



# The Mount Logan (Yukon) ice cores:

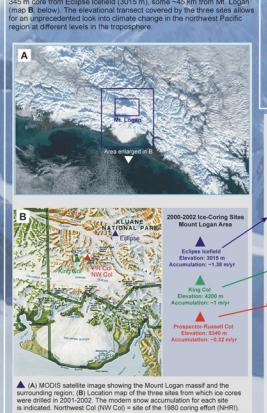
## A time window into northwest Pacific climate variability



### Principal Investigators and their affiliation:

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**Introduction**  
 In the summers of 2001 and 2002, scientists from the Geological Survey of Canada (GSC), Japan's National Institute for Polar Research (NIPR), the University of New Hampshire (UNH), and University of Maine (UMaine) recovered three ice cores from Mount Logan (elevation 3,215 m above sea level; elevation 5,920 m a.s.l.) in the St. Elias Mountain range of south-western Yukon (map A, below). The GSC recovered a 167 m core to bedrock from Mount Logan (Fig. 1A), while UNH recovered a ~58 m core from Mt. Logan (map B, below). The area of interest covered by the three sites allows for comparison to locations close to climate change in the northwest Pacific region at different levels in the troposphere.



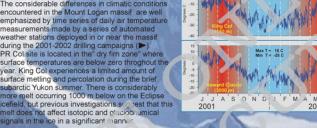
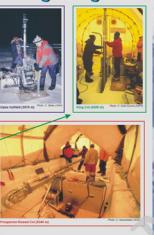
### Vertical contrast

Previous work on ice cores from the summit of Mt. Logan by G. Holdsworth and colleagues (Env. Canada/NHRI Expedition, 1980) indicate that the PR Col core should provide a record of temperature, snow accumulation, atmospheric composition and circulation patterns in the atmosphere. A preliminary analysis of the ice core data indicates that the climatic signals recorded in the Mt. Logan core before the 1980 coring project, indicate that the climate record at this site covers the entire Holocene period (c. 11,500 years), with an expected temporal resolution varying from annual (top of core) to centennial (bottom).

Glaciological records developed from a 150 m core recovered from Eclipse icefield in 1998 by O. Goto-Azuma (NIPR) and colleagues (Env. Canada/NHRI Expedition, 1998) indicate the record of atmospheric pollution and volcanic eruptions (acidic fallout), as well as closer connections with surface temperature and pressure records from regional weather stations. Because it was drilled from an area of low accumulation, the Eclipse icefield core is shorter than the Mt. Logan core (c. 345 m), but higher-resolution climate record than the Mt. Logan core (c. 1,500 years) should be provided by the Eclipse icefield core.

Because the three cores span an elevation range of over 2,000 m, they will provide a unique opportunity to investigate changes in the vertical structure of the atmosphere in relation to climate change, and also to test the reliability of the Holocene periods. Analyses of the three cores will be initiated in 2003 from Mount Logan to Churchill in nearby Manitoba by team from the Byrd Polar Research Centre (Ohio State University).

### Drilling through time

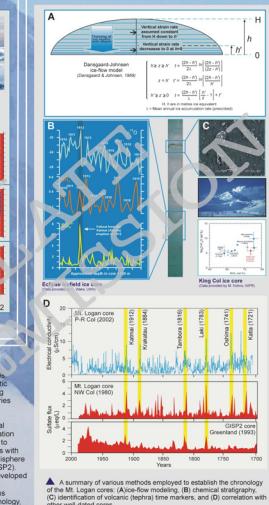


The considerable differences in climatic conditions encountered in the Mount Logan massif are well emphasized by the time series of daily air temperature and daily air density measured at the three atmospheric weather stations deployed in or near the massif (see Fig. 1B). The Mount Logan atmospheric station (PR Col) is located in the "dry fir zone" where similar King Mettler temperatures are recorded as are seen at King Col, except during a short amount of surface melting and percolation during the brief summer months (May-July). The daily air density at PR Col is occurring 1,000 m below on the Eclipse icefield; previous investigations have tested that this small difference in elevation does not affect the PR Col ice core air density values in a significant manner.

### Dating the Mt. Logan ice cores

The most challenging thing up in building reliable chronologies for Mount Logan ice cores. This is an essential, and often time-consuming process that must be taken before it can define, and interpret the ice cores. The chronology can be formulated. Following the interval 2001-2002, using three methods were employed to establish the chronology of the three cores. These methods involved the USA, Denmark and Japan. A combination of methods were used to determine the chronology of the cores. The methods include: (A) the use of  $^{210}\text{Po}$  in the upper layer; (B) the use of  $^{14}\text{C}$  in the upper layer; (C) the identification of annual layers from the presence and absence of annual layering; and (D) cross-correlations with other well-dated cores. (A)  $^{210}\text{Po}$  dating is used to date the upper layers due to the short time scale of use (years). This method is to date the upper layers due to source eruptions of known date, and (D) cross-correlations with other well-dated cores. (B) the use of  $^{14}\text{C}$  (through a geochronological model); (C) the identification of annual layers due to source eruptions of known date, and (D) cross-correlations with other well-dated cores. (A) A preliminary scale for the PR Col core has been developed for the PR Col core. This scale is based on the  $^{14}\text{C}$  data (and similar scales for  $^{210}\text{Po}$  for acidity). The initial interpretations discussed here are based on this preliminary chronology, and should therefore be considered speculative for the time being.

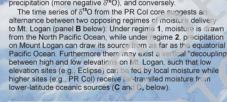
Slicing an ice core from Mt. Logan.



A summary of various methods employed to establish the chronology of the Mount Logan ice cores. (A) Schematic diagram of the Mount Logan ice cores. (B) Time series plot of air temperature (°C). (C) Time series plot of air density ( $\text{kg m}^{-3}$ ). (D) Time series plot of air density ( $\text{kg m}^{-3}$ ) compared with other well-dated cores.

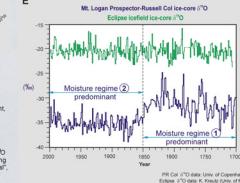
### Isotopes in precipitation: Clues to climate variability

The ratio of stable O and H isotopes in precipitation is determined by several factors, including air temperature, distance from the oceans, and source of water vapor, as well as the ratio of sea surface temperature (SST), net evaporation, and precipitation. In polar regions the ratio of oxygen isotopes in snow (expressed by  $\delta^{18}\text{O}$ ) as a function of time can be used to track climate change induced by the temperature at the time of precipitation. On Mount Logan however, we hypothesize that distance to the moisture source is much more important than air temperature. On Mount Logan however, the isotopic "fingerprint" follows the trajectory of the air mass that carried the precipitation to the site of precipitation. We expect that the isotopic signature of precipitation will change as the air mass moves away from the source region, with the highest  $\delta^{18}\text{O}$  values occurring at high elevations on Mt. Logan, such that low elevation ice cores reflect the isotopic signature of precipitation from lower-latitude oceanic sources (C and E, below).

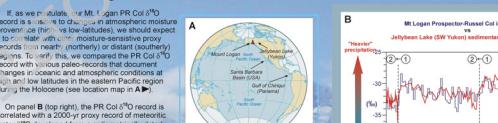


### Abrupt Climatic shifts in Mt. Logan ice?

We previously examined the Mt. Logan and P.R. Col ice cores to evaluate the possibility of abrupt shifts in the Holocene period. We used "abrupt shifts," which are punctuated events, during which the isotope signal shifts from one contrasting mode to another. These shifts are observed on many ice cores, and are interpreted as reflecting large-scale atmospheric circulation changes. In the case of the two ice cores from Mt. Logan, there is no sign of such a shift in either the Mt. Logan or the Eclipse icecore, despite some differences in the precipitation regimes on the two icecores. Thus, we find no evidence for abrupt shifts in the Mt. Logan ice cores. However, we find some evidence for abrupt shifts in the P.R. Col ice core, despite some differences in the precipitation regimes on the two icecores. This finding raises the possibility that future warming scenarios may stimulate a natural, built-in tendency for climate to shift on its own when some threshold condition is exceeded.



### The Mount Logan ice-core record and Pacific climate teleconnections



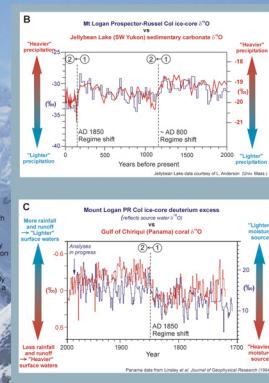
On panel B (top right), we compare the PR Col ice core with a 2000-yr proxy record of meteoric water from Jellybean Lake in the southern Yukon. The two cores are close in age (within 15 years) and are closely correlated in both as an earlier shift at AD 800 possibly related to the "Regime shift" in the PR Col core at AD 1650. This comparison suggests that the "Heavier" precipitation regime documented in the Mt. Logan icecore "sees" the precipitation regime documented in the Jellybean Lake icecore, possibly because of the strongly constraining  $\delta^{18}\text{O}$  signatures in distant or proximal precipitation.

On panel C (middle right), we compare changes in relative humidity of water vapor in precipitation with the PR Col ice core, with a record of SST variations measured in growth bands of a coralline alga in the Gulf of California. On the production of Pacific SSTs, the correlation is not as strong as the SST correlation with the Mt. Logan ice core. This is possibly because the  $\delta^{18}\text{O}$  values in China corals are more strongly affected by the isotopic composition of the surface ocean rather than by rainfall or evaporation conditions, which influences the SSTs. However, changes in SSTs may affect the  $\delta^{18}\text{O}$  signature of ice cores growth layers toward lower (more negative)  $\delta^{18}\text{O}$  values. Thus, while SSTs may be a useful proxy for the hydrological balance in the coastal (Pacific) region of Panama, there again we find a reasonably good agreement with the PR Col ice core, with the AD 1650 regime shift clearly evident in the SSTs. This regime shift in SSTs is well documented by the shift in coral growth layer thickness on Mt. Logan and may be a useful proxy for the hydrological balance in the coastal (Pacific) region of Panama.

Again one would point the frequency, persistence, magnitude and location of the  $\delta^{18}\text{O}$  shifts throughout the Holocene period. The regime shift (5 m in length) between 1600 and 5000 years ago is sampled at the University of Maine at much higher resolution for the PR Col and Gulf of California analyses. This will allow us to further investigate the suddenness of the regime shifts observed in the preliminary record.

It appears from these comparisons that the Pacific regime shifts are more climatically connected under moisture regime 2 periods. In an earlier conceptual framework, one would say that during climate regime 1 the Mt. Logan ice core receives more precipitation from a single moisture source feeding precipitation in Mt. Logan to the relatively dry North Pacific waters (moisture regime 1 situation). The new precipitation pattern involves a mixing of two moisture sources (moisture regime 2 situation). D. Dahlman et al. (unpublished) will be of extreme interest in the Mt. Logan ice core. We anticipate that as new results emerge from the Mt. Logan ice core, we will further elucidate the Holocene climate history of the Northwest Pacific region.

For further reading:  
 Fidell, J. E. et al. (2003) Geophysical Research Letters v. 30, no. 11  
 Holdsworth, G. J. et al. (1980) Journal of Glaciology v. 26, pp. 57-63  
 Holdsworth, G. J. et al. (1989) Journal of Geophysical Research v. 94, pp. 1483-1498.  
 Holdsworth, G. J. et al. (1998) Journal of Climate v. 11, pp. 1833-1844.  
 Linsley, R. K. et al. (1989) Journal of Climate v. 2, pp. 979-988.  
 Moore, G.W.K. et al. (2002) Annales of Glaciology v. 35, pp. 413-429.  
 Wake, C.P. et al. (2002) Annales of Glaciology v. 36, pp. 416-422.



Comparisons between the stable O isotope record of the Mt. Logan (blue) and Santa Barbara (green) ice cores. Note the different time scales for each panel. Colored arrows provide a rough estimate of the time scale for each core. The arrows for the ice cores represent the time scale for the records after the core was coring. See main text for discussion.